

### STOP 1: Lower Gauja spillway valley at Sigulda

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The stop (24°50'2"E, 57°10'15"N, for location see Fig. 1.1) at the Krimulda Castle ruins in Sigulda introduces the geological structure, morphology and formation of the River Gauja valley and slope processes between the towns of Valmiera and Vangaži. This stretch of the river valley, also known as the Lower Gauja spillway valley, is about 110 km long. The spillway is confined to an ancient buried valley incised into Middle and Upper Devonian sedimentary rock (Pērkons 1947). The oversized river valley is 1-2.5 km wide and reaches a depth of 25 m near Valmiera, 35-40 m in the vicinity of Cēsis and 85 m at Sigulda. The floor of the bedrock within the valley lies at 17-18 m a.s.l. near Valmiera, 12 m b.s.l. in the vicinity of Cēsis and more than 50 m b.s.l. at Sigulda (Fig. 1.2). It is significant that downstream of Cēsis the ancient valley is carved into weakly-cemented and/or soft Devonian terrigenous rock that runs along the lithological boundary with carbonate rock.

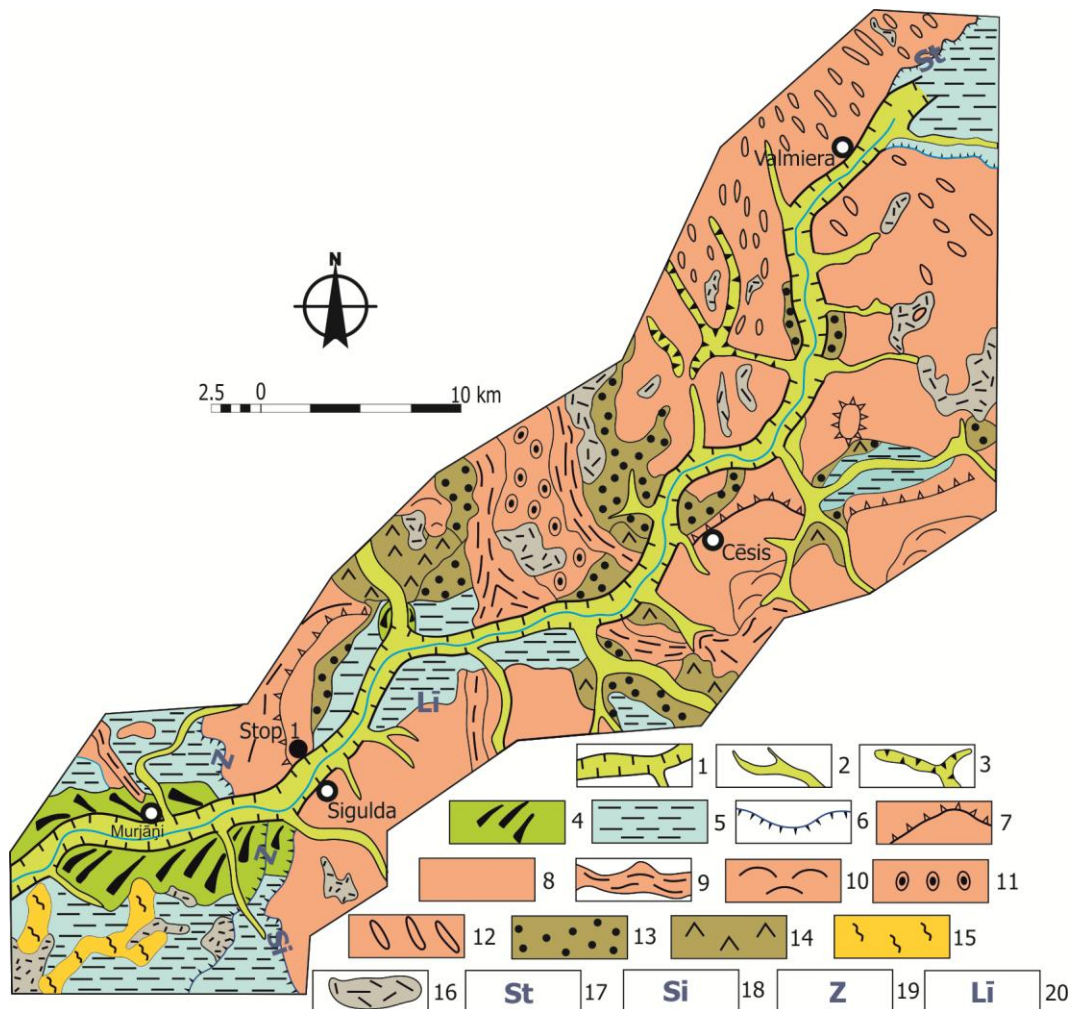


Fig. 1.1. Geomorphological map of the River Gauja valley and the adjacent area between Valmiera town and Murjāņi village. 1 – Lower Gauja spillway valley; 2 – valleys of tributaries; 3 – largest gullies; 4 – late-glacial delta plains; 5 – glaciolacustrine plains; 6 – ancient shorelines of glacial lakes; 7 – ice-contact and bedrock scarps; 8 – till plains; 9 – ice marginal ridges; 10 – morainic hills; 11 – cupola-like hills; 12 – drumlins; 13 – glaciofluvial plains; 14 – kames; 15 – inland dunes; 16 – mires; 17 – Strenči proglacial lake; 18 – Silciems ice-dammed lake; 19 – Zemgale ice-dammed lake; 20 – Līgatne ice-dammed lake; black circle – location of the stop.

According to Āboltiņš (1971), the buried valley was formed before the Saalian glaciation. This conclusion is based on the occurrence of the lower till bed on the surface of the Middle Devonian sandstone at the base of the buried valley near the towns of Valmiera, Cēsis and Sigulda, and Murjāņi village (Figs. 1.2 and 1.3). The cross-section of the Lower Gauja spillway valley is asymmetrical, with a general prevalence of erosional terraces. These terraces represent the Sigulda terrace spectrum of the Gauja River valley (Āboltiņš 1971).

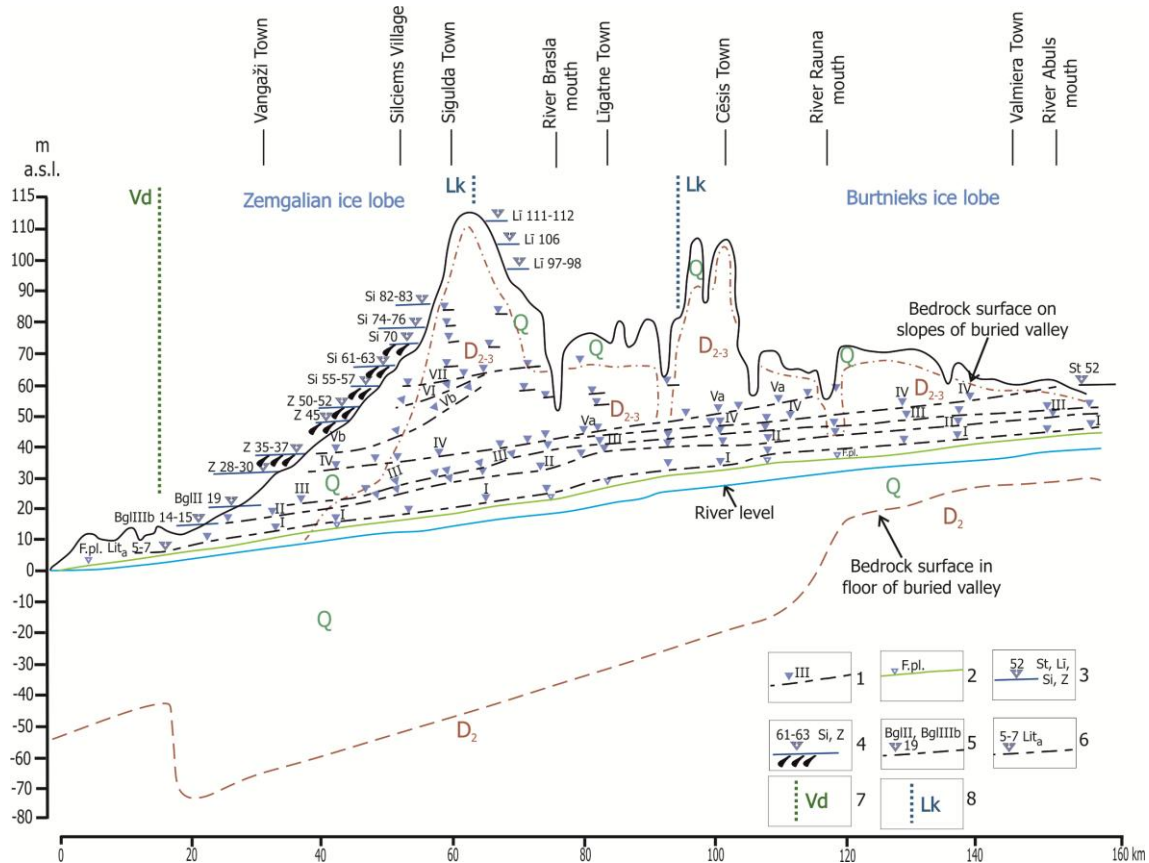


Fig. 1.2. Longitudinal profile of valley and terraces of the River Gauja downstream of the Strenči meltwater basin. 1 – terrace and its symbol; 2 – floodplain; 3 – shoreline elevation and symbols of glacial lakes: St – Strenči meltwater basin, Lī – Līgatne ice-dammed lake; Si – Silciems ice-dammed lake, Z – Zemgale ice-dammed lake; 4 – elevations of late-glacial delta levels in Silciems and Zemgale ice-dammed lakes; 5 – elevation and symbol of the Baltic Ice Lake stage Bgl II and phase Bgl III<sub>b</sub>; 6 – elevation and symbol of the Littorina Sea maximum transgression phase (Lit<sub>a</sub>); 7 – ice marginal position of the Valdemārpils phase; 8 – ice marginal position of the Linkuva phase.

On the basis of geological and geomorphological studies Āboltiņš (1971) distinguished seven terrace levels in the Gauja River valley at Sigulda, but some terrace remnants are located on the upper part of the valley slopes upstream and downstream of this town, at a higher elevation than terrace VII (Fig. 1.3). According to Āboltiņš (1969, 1971), the highest terraces of the upper complex (terraces VI and VII) were formed by meltwater streams, which flowed from melting dead ice and small proglacial lakes, located in areas near the valley, into the Silciems ice-dammed lake. Terraces IV and V were produced as a result of water drainage from Strenči meltwater basin, located in the topographically lowest part the Northern Vidzeme Lowland, into the Zemgale ice-dammed lake, which partly occupied the lower part of the Central Latvian Lowland.

The latest studies of the Līgatne ice-dammed lake terraces between Sigulda and the River Amata testify that the highest terraces of the upper complex of the River Gauja

represent the shorelines of an ice meltwater basin. Terraces developed during several phases of incision are related to the evolution of the Līgatne ice-dammed lake. Next to the Ratnieki Conference and Recreational Centre of the University of Latvia, situated on a glaciolacustrine plain, the terraces of the Līgatne ice-dammed lake can be traced at 111-112 m, 106 m and 97-98 m a.s.l. (Fig. 1.2). Terraces composed of fine-grained sand, silt and clay are traceable at levels of 75, 61-62 and 51-52 m a.s.l. The thickness of the glaciolacustrine sediments varies from some tens of centimetres to a few metres. The River Gauja terraces formed by streams are traceable at levels below 50 m a.s.l.

Both highest terraces of the lower complex (terraces III and II) are related to levels of the Baltic Ice Lake (stage Bgl II and phase Bgl III<sub>b</sub>), whereas terrace I conjugates with the Littorina Sea phase Lit<sub>a</sub> level (Āboltniš 1971; Grīnbergs 1957).

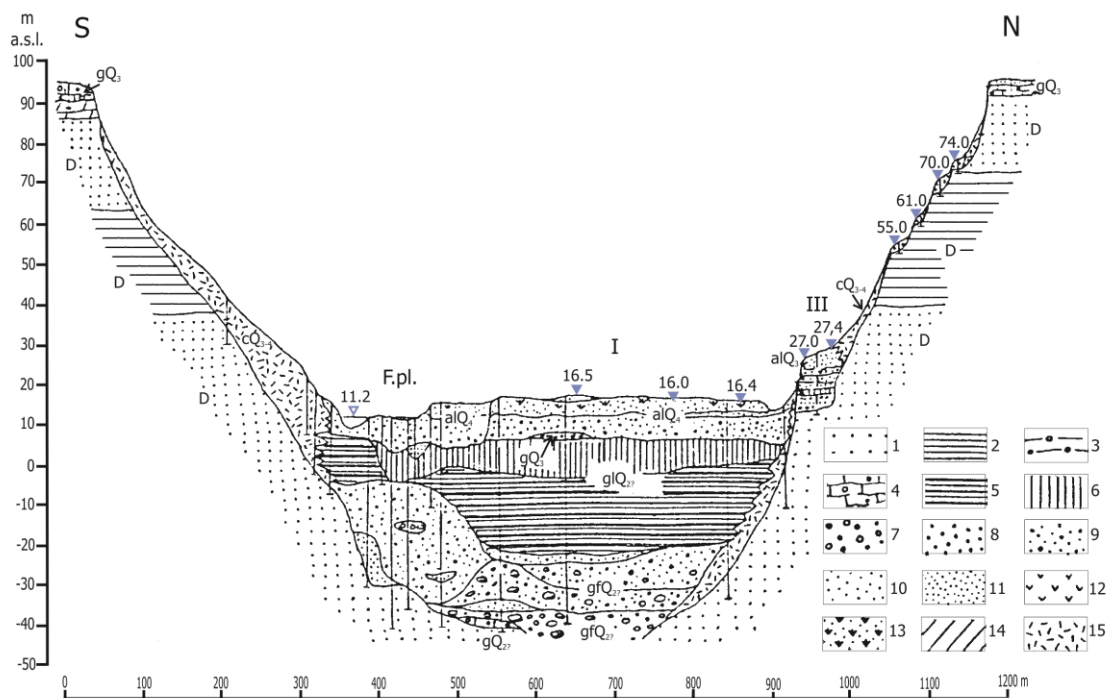


Fig. 1.3. Cross-section of the ancient River Gauja valley at the town of Sigulda (modified after Āboltniš 1971 and Pērkonis 1947). 1 – Devonian sandstone; 2 – Devonian clay and siltstone; 3 – Saalian (?) till (gQ2); 4 – Weichselian reddish brown sandy-clayey basal till (gQ3); 5 – varved clay (glQ2?); 6 – silt (glQ3 and alQ4); 7 – pebbles with gravel (alQ3 and alQ4); 8 – gravel (alQ3 and alQ4); 9 – sand with gravel (alQ3 and alQ4); 10 – fine-grained sand (alQ3 or alQ4); 11 – silty fine-grained sand (alQ3 or alQ4); 12 – peat (pQ4); 13 – disseminated organic matter in sand (alQ4); 14 – silty loam (alQ4); 15 – sandy colluvium (cQ3-4).

Formation of the stretch of the River Gauja between the towns of Valmiera and Vangaži began after the ice retreat from the marginal zone of the Linkuva phase (North Lithuanian) at least by about 15.2 cal. ka BP (calibrated from 13.0 ka BP using the IntCal09 calibration curve). Terraces VII to IV were apparently formed before the Allerød. According to Stelle et al. (1975 a, b), the radiocarbon age obtained from organic debris at the Viesulēni farmhouse is  $11,270 \pm 230$   $^{14}\text{C}$  yrs BP (Ri-105) and  $11,114 \pm 350$   $^{14}\text{C}$  yrs BP (Ri-74), and peat sediments at the Līči Sanatorium have been dated to  $10,535 \pm 250$   $^{14}\text{C}$  yrs BP (Ri-33) and  $10,282 \pm 250$   $^{14}\text{C}$  yrs BP (Ri-33A). These data refer to the formation of terraces III and II during the Allerød and Younger Dryas. Terrace I developed during a relatively long time interval of the Holocene (from the Boreal until the Subboreal). All the terraces of the uppermost complex have been generated under conditions of continuous intensive downcutting. The alluvium of the lower terrace complex and floodplain consists of two members, i.e. includes both channel and

floodplain facies. The thickness of the alluvium reaches up to 8-10 m and more in places. The lowest part of the Gauja valley is occupied by an aggradational floodplain (Figs. 1.2 and 1.3).

The slopes of the Lower Gauja spillway valley are also dissected by numerous gullies and valleys of tributaries, particularly in the stretch between the towns of Līgatne and Sigulda. The slopes of the River Gauja and its tributaries display numerous landslides, suffosion caves and grottoes, along with typical and valley-like gullies. Alluvial-proluvial fans, and cave and grotto floors are related to different terrace levels, which suggests successive formation of these landforms in the course of development of the River Gauja valley.

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LATVIAN ASSOCIATION FOR QUATERNARY RESEARCH

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Organized by:

University of Latvia  
Daugavpils University  
Latvian Association for Quaternary Research  
INQUA Peribaltic Working Group (INQUA TERPRO Commission)

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Editors: Vitālijs Zelčs and Māris Nartišs

The English texts of the field guide were revised by Valdis Bērziņš

Recommended reference for this publication:

Zelčs, V. and Nartišs, M. (eds.) 2014. Late Quaternary terrestrial processes, sediments and history: from glacial to postglacial environments. Excursion guide and abstracts of the INQUA Peribaltic Working Group Meeting and field excursion in Eastern and Central Latvia, August 17-22, 2014. University of Latvia, Rīga, 2014, 150 pages.

Sponsored by:

University of Latvia

Layout: Vitālijs Zelčs, Māris Nartišs and Māris Krievāns

ISBN 078-9934-517-60-0

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This volume is available from:

Faculty of Geography and Earth Sciences  
University of Latvia  
Rainis Blvd. 19  
Rīga, LV1586  
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